

Principles of radioactive decay

- Clues from isotopes

- "Chemical element"
specified by # of electrons in
the neutral atom
= atomic number (Z)
- Example: Rubidium (Rb) has 37
electrons. $Z=37$

Review : Isotopes

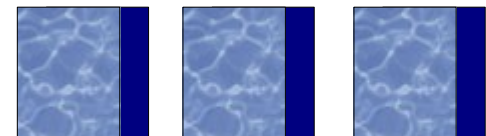
□ Isotopes are atoms that have the same atomic number Z but difference **atomic weights (A)**

□ Review: Hydrogen has 3 isotopes

1) H^1 Hydrogen 1 p 0 n 1e

2) H^2 Deuterium 1 p 1 n 1e

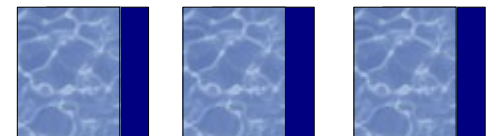
3) H^3 Tritium 1 p 2 n 1e



Example two: Rubidium

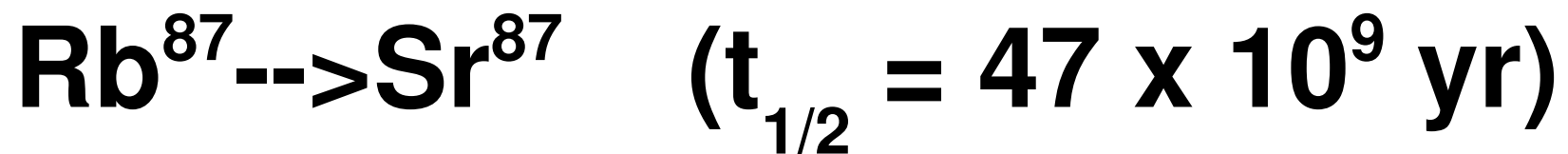
- Rubidium has 2 common isotopes

	# e	#p	#n
A	Z	Z	(A-Z)
Rb⁸⁵	37(-)	37(+)	48
Rb⁸⁷	37(-)	37(+)	50

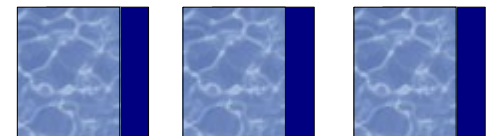


Rubidium 87 is **radioactive**

- Decays spontaneously into Sr^{87} (strontium) with a **half-life** of 47 billion years:

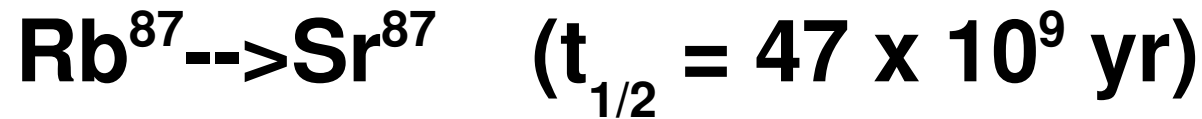


What does half-life mean?



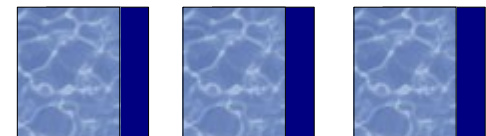
Radioactive half-life

- Decays spontaneously into Sr^{87} (strontium) with a **half-life** of 47 billion years:

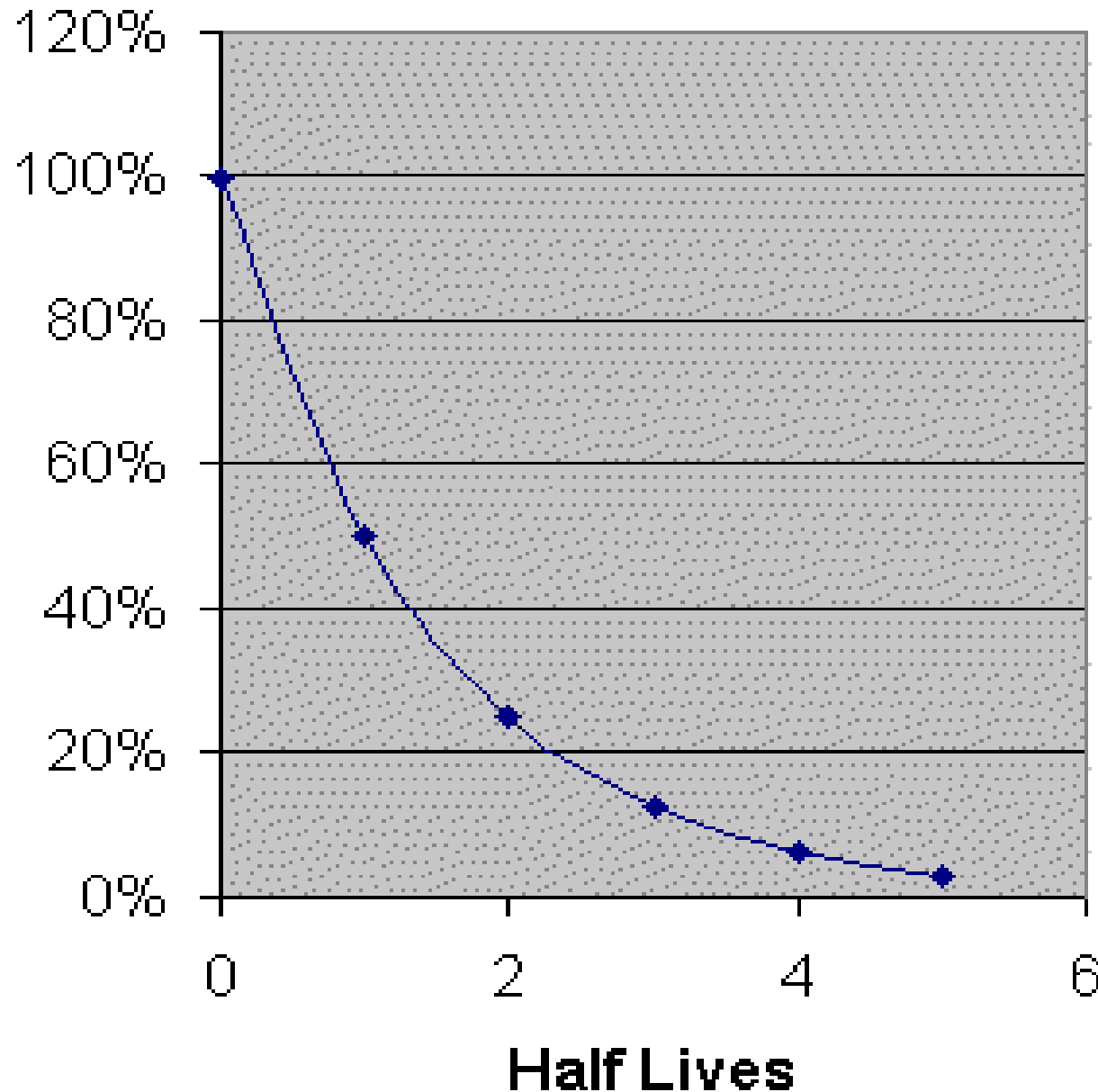


What does half-life mean?

If we start with N atoms of Rb^{87} at time $t=0$, after 47 Gyr there will be $N/2$ left.



Radioactive Decay Curve

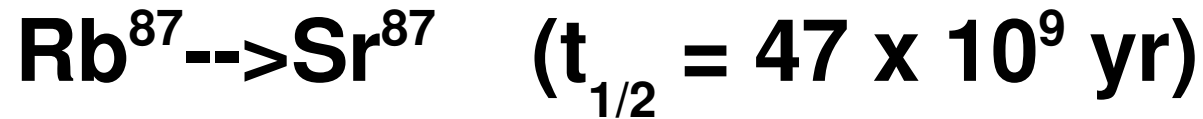


◆ Parent Isotope Remaining

With each half-life, half the remaining radioactive element disappears.

Parent-daughter

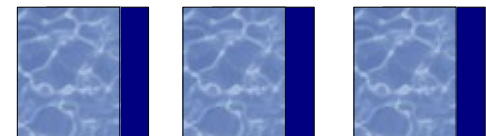
□ In this decay:



Rb^{87} is called the *parent* isotope

Sr^{87} is called the *daughter* isotope

The daughter isotope abundance *increases* with time.



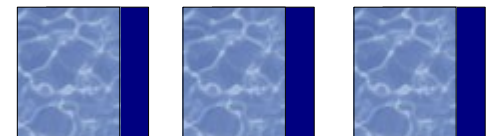
Chronology: minerals

$\text{Rb}^{87} \rightarrow \text{Sr}^{87}$ (unstable)

Sr^{87} is STABLE (doesn't decay).

How can we use this?

When rocks condense, they contain specific minerals, some of which incorporate Rb and Sr : this process ignores the isotope. i.e., will incorporate as much Rb^{87} (unstable) as Rb^{86} (stable).



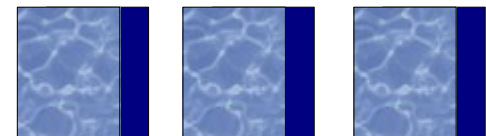
Subsequent decay

- ❑ So some minerals contain more Rb than others.
- ❑ The Rb^{87} decays to Sr^{87}
- ❑ If we could measure Sr^{87} in one of these minerals , could we know how long ago the rock formed?
- ❑ NO: for two reasons



2 practical problems

- ❑ There may often be some initial daughter element!
- ❑ Measurements of abundance of a single isotope very difficult. But one can measure RATIOS very very accurately.



The trick to absolute isotopic cosmochronology

(besides spelling it!)

- Divide by the abundance of another isotope of the daughter element:

$\text{Rb}^{87} \rightarrow \text{Sr}^{87}$; use Sr^{86} as reference

So measure the ratios :

$\text{Sr}^{87}/\text{Sr}^{86}$ and $\text{Rb}^{87}/\text{Sr}^{86}$

with mass spectroscopy

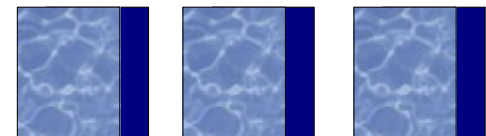
(transparency)



Isotopic evolution from radioactive decay produces an **isochrone**

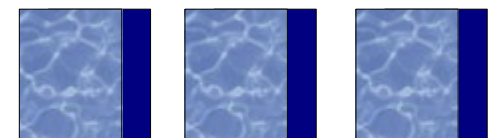
- **Slope of plot of $(\text{Sr}^{87}/\text{Sr}^{86})$ versus $(\text{Rb}^{87}/\text{Sr}^{86})$ gives (see transparency):**
 - **Time since rock solidified**
 - **If system has been 'closed'**
 - **Initial abundance of daughter element (see transparency)**

Have to measure several minerals.



Rb/Sr dating of meteorites

- All primitive chondrites give ages of 4.5 +/- 0.1 Gyr.
 - Some of the best are 4.54 +/- 0.05 Gyr
- Contrast: Oldest rocks on Earth about 3 Gyr old (why? Is Earth 3 Gyr old?)
- Oldest rocks on moon ~4.3 Gyr old
 - Chondrites date the 'age of the Solar System'

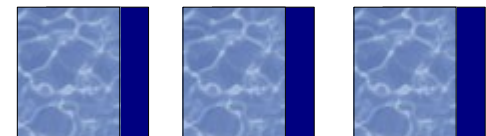


I/Xe isotopic system

- Iodine 129 decays spontaneously (into Xenon 129) with a **half-life** of 17 million years:



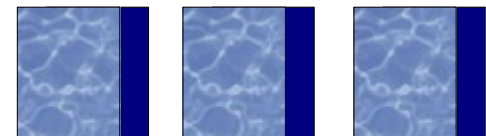
But Xe is a noble gas, so it wouldn't be normally found in rocks???



I/Xe isotopic system

- ❑ **Solution: Iodine must have been incorporated into some rocks and then decayed into Xe. The Xe could not escape ('trapped') IF the rock was cool (<600 K).**
- ❑ **Consequence: The Xe¹²⁹ can be measured and tell us how long it took the 'parent body' to cool.**

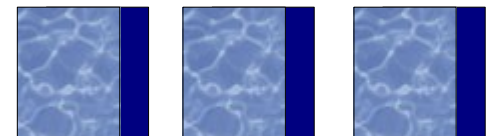
What do measurements say?



Meteoritic 'Parent Body' cooling times

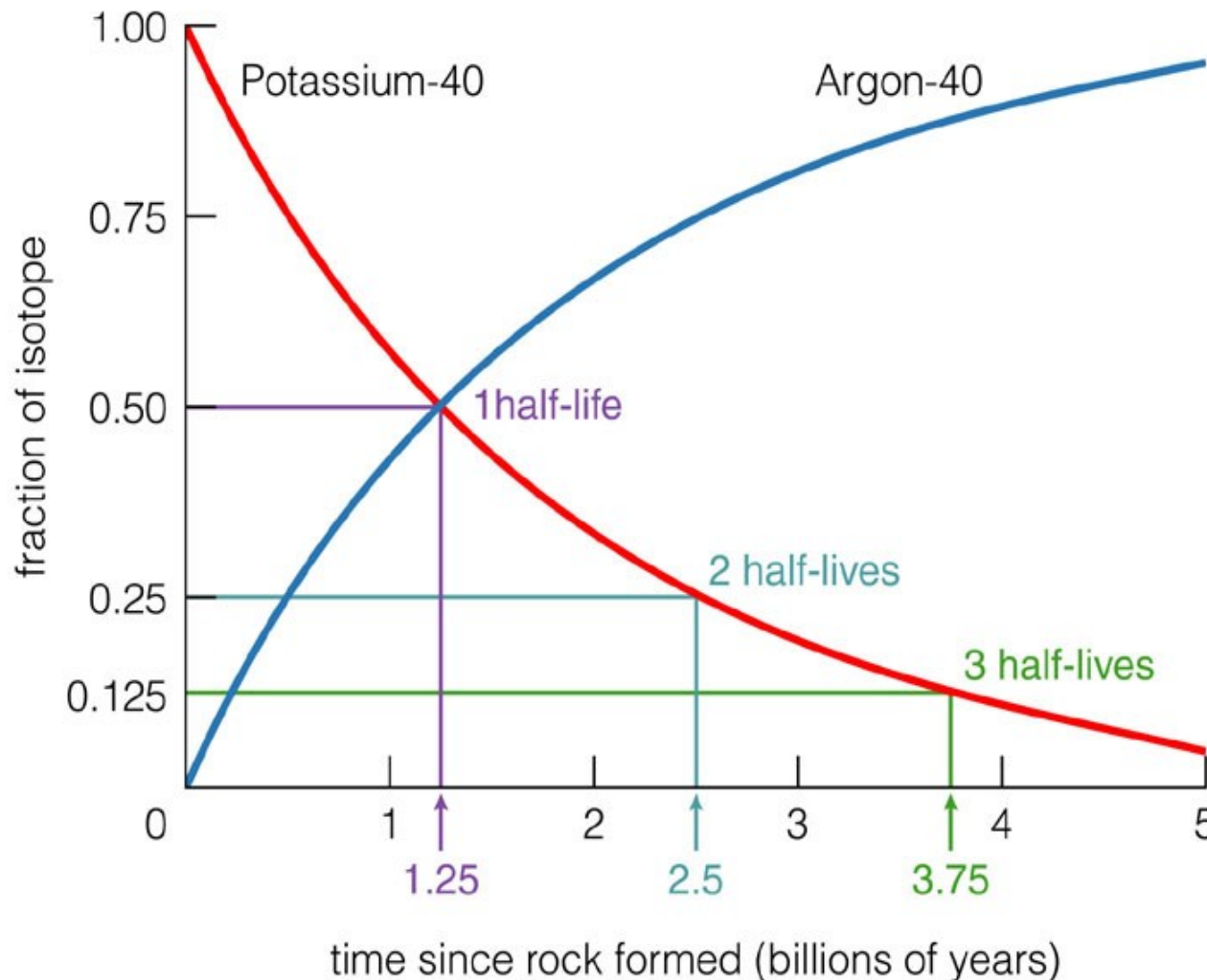
- ❑ Answer: Parent bodies typically took 10-100 Myr to cool
 - ❑ Most chondrites: ~20 Myr
 - ❑ Achondrites/irons: ~100 Myr
- ❑ So differentiated parent bodies (measured from achondrite stones) took longer to cool.
Reasonable!

Important implication: Differentiated PBs formed (accretion + melting + cooling) in <100 Myr and *never reheated* (why?)



K-Ar system

- ❑ Potassium 40 decays into Argon 40
- ❑ Argon 40 escapes from warm rocks



Presence of Argon 40 dates the 'cooling age' of the rock.

Some lunar rocks 4.4 Gyr old --> Moon formed early.

